**Effects of in-situ and reanalysis climate data on estimation of cropland gross primary
production using the Vegetation Photosynthesis Model**

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Abstract: Satellite-based Production Efficiency Models (PEMs) often require meteorological reanalysis data such as the North America Regional Reanalysis (NARR) by the National Centers for Environmental Prediction (NCEP) as model inputs to simulate gross primary production (GPP) at regional and global scales. This study first evaluated the accuracies of air temperature (T_{NARR}) and downward shortwave radiation (R_{NARR}) of the NARR by comparing with in-situ meteorological measurements at 37 AmeriFlux non-crop eddy flux sites, then used one PEM—the Vegetation Photosynthesis Model (VPM) to simulate 8-day mean GPP (GPP_{VPM}) at seven AmeriFlux crop sites, and investigated the uncertainties in GPP_{VPM} from climate inputs as compared with eddy covariance-based GPP (GPP_{EC}). Results showed that T_{NARR} agreed well with in-situ measurements; R_{NARR} , however, was positively biased. An empirical linear correction was applied to R_{NARR} , and significantly reduced the relative error of R_{NARR} by ~25 % for crop site-years. Overall, GPP_{VPM} calculated from the in-situ ($\text{GPP}_{\text{VPM}(\text{EC})}$), original ($\text{GPP}_{\text{VPM}(\text{NARR})}$) and adjusted NARR ($\text{GPP}_{\text{VPM}(\text{adjNARR})}$) climate data tracked the seasonality of GPP_{EC} well, albeit with different degrees of biases. $\text{GPP}_{\text{VPM}(\text{EC})}$ showed a good match with GPP_{EC} for maize (*Zea mays* L.), but was slightly underestimated for soybean (*Glycine max* L.). Replacing the in-situ climate data with the NARR resulted in a significant overestimation of $\text{GPP}_{\text{VPM}(\text{NARR})}$ (18.4/29.6 % for irrigated/rainfed maize and 12.7/12.5 % for irrigated/rainfed soybean). $\text{GPP}_{\text{VPM}(\text{adjNARR})}$ showed a good agreement with $\text{GPP}_{\text{VPM}(\text{EC})}$ for both crops due to the reduction in the bias of R_{NARR} . The results imply that the bias of R_{NARR} introduced significant uncertainties into the PEM-based GPP estimates, suggesting that more accurate surface radiation datasets are needed to estimate primary production of terrestrial ecosystems at regional and global scales.

Keywords: Vegetation Photosynthesis Model (VPM), NARR, MODIS, AmeriFlux, downward shortwave radiation, vegetation indices

1. Introduction

Croplands cover 12 % of the global ice-free terrestrial surface (Ramankutty et al., 2008) and provide food for more than seven billion people in the world. Increasing demand for food under the changing climate is one of the great challenges in the coming decades (Guanter et al., 2014; Lobell and Asner, 2003). **Gross Primary Production (GPP)** of croplands is the total carbon uptake through photosynthesis. A recent modeling study estimated that croplands have an annual sum of 11 Pg C yr⁻¹ GPP, accounting for ~10 % of the global terrestrial GPP (Chen et al., 2014). Crop cultivation and production vary substantially over space and time. **Thus**, an accurate quantification of cropland GPP is critical for global food security (Wheeler and von Braun, 2013), biofuel production (Landis et al., 2008), and understanding variations in the terrestrial carbon cycle (Haberl et al., 2007).

Production Efficiency Models (PEMs) have been widely used to quantify the spatial-temporal GPP variations of terrestrial ecosystems using the satellite and climate data as inputs. The PEMs, originating from Monteith's theoretical concept about light use efficiency (LUE) (Monteith, 1972; Monteith and Moss, 1977), estimate GPP as the product of the photosynthetically active radiation (PAR, MJ m⁻²), the fraction of PAR absorbed by the vegetation (fPAR), and the conversion efficiency of absorbed PAR for carbon fixation (ϵ , g C MJ⁻¹) ($GPP = \epsilon \times fPAR \times PAR$). The PEMs for croplands can be classified into two categories based on fPAR and ϵ estimation methods. The first category calculates fPAR and ϵ separately. This approach has been applied in the Global Production Efficiency Model (GLO-PEM) (Prince

and Goward, 1995), the MODIS Daily Photosynthesis model (MODIS-PSN) (Running et al., 2000), the C-Fix model (Veroustraete et al., 2002), and the Vegetation Photosynthesis Model (VPM) (Xiao et al., 2004a; Xiao et al., 2004b). The second type of PEMs, referred as the Greenness and Radiation (GR) model, uses the chlorophyll-related vegetation indices (VI_{chl}) as a proxy of $\epsilon \times fPAR$ ($GPP \propto VI_{chl} \times PAR$) (Gitelson et al., 2006; Peng and Gitelson, 2011; Peng and Gitelson, 2012; Peng et al., 2011; Wu et al., 2009; Zhang et al., 2014; Zhang et al., 2015).

Challenges remain, however, in applying PEMs due to model structure and model inputs. Several attempts have been made to address the uncertainties from the PEM algorithm itself, including the assumption of linear response of photosynthesis to light intensity (Chen et al., 1999), constant maximum LUE for one ecosystem (Heinsch et al., 2006), the impacts of diffuse radiation (He et al., 2013; Zhang et al., 2012), and the incomplete integration of environmental regulations (temperature, water, phenology etc.) to photosynthetic processes (Dong et al., 2015; Yuan et al., 2014). Most uncertainty analyses overlooked the potential impacts of model inputs on the application of PEMs to regional or global primary production monitoring.

Meteorological reanalysis data produces continuous and near real-time climate monitoring via data assimilation models, and has been the major climate input of PEMs for the large-scale primary production simulation (Feng et al., 2007; Running et al., 2004; Xiao et al., 2011; Yuan et al., 2010). Studies have reported that the meteorological reanalysis data can be spatially and temporally biased from the ground observations, in particular for downward shortwave radiation when estimating PAR (Babst et al., 2008; Cai et al., 2014; Decker et al., 2012; Troy and Wood, 2009; Zhang et al., 2007; Zhao et al., 2013a; Zhao et al., 2006; Zib et al., 2012). PEMs have been found very sensitive to the accuracy of climate reanalysis variables (Cai et al., 2014; Heinsch et al., 2006; Zhang et al., 2007; Zhao et al., 2006). For example, Heinsch et

al. (2006) reported that the errors associated with the standard MODIS GPP product were mainly attributed to the NASA's Data Assimilation Office (DAO) reanalysis data. Previous sensitivity analyses of PEMs to climate inputs focused on global reanalysis data, the spatial resolution of which is too coarse to delineate the local climatic variations.

The North America Regional Reanalysis (NARR) by the National Centers for Environmental Prediction (NCEP) is the only currently available long-term regional reanalysis data. Compared with the NCEP global reanalysis datasets, the NARR substantially improves the spatio-temporal resolutions along with the accuracy of climate variables (Mesinger et al., 2006) and could be an alternative climate driver of regional GPP estimates in particular for croplands, one of the most heterogeneous landscapes. There has been very limited research regarding the uncertainties of PEMs in relation to the NARR. Therefore, careful investigation of the accuracy of the NARR and its impacts on cropland GPP estimates at site level is an indispensable step prior to the large scale application of these tools.

The objectives of this study were to: (1) evaluate the accuracy of the NARR (air temperature and downward shortwave radiation) as compared to the in-situ observations from the AmeriFlux network at 8-day intervals; (2) adjust the NARR based on the statistical differences from in-situ meteorological measurements; and (3) quantify the impacts of different climate inputs (in-situ meteorological data and the original and adjusted NARR data) on the GPP simulation for maize and soybean using the VPM at seven AmeriFlux crop sites (40 site-years).

2. Data and Methods

2.1. NARR

The NARR is produced at a spatial resolution of 32 km and a temporal resolution of 3-hours. We obtained the NARR daily gridded air temperature (T_{NARR}) and downward shortwave radiation (R_{NARR}) from <http://www.esrl.noaa.gov/psd/>. The daily T_{NARR} and R_{NARR} for the pixels covering AmeriFlux sites were extracted for the available site-years at 44 AmeriFlux sites and were aggregated to 8-day intervals to match the temporal resolution of MODIS products.

2.3. MODIS land surface reflectance, vegetation indices products

This study used the 8-day 500 m MODIS Surface Reflectance product—MOD09A1 to derive vegetation indices. The time-series MOD09A1 data for the crop sites were extracted from the MODIS data portal at the Earth Observation and Modeling Facility (EOMF), University of Oklahoma (<http://www.eomf.ou.edu/visualization/manual/>). The Enhanced Vegetation Index (EVI) and Land Surface Water Index (LSWI) were calculated for every 8-day observation using equations (1) and (2).

$$EVI = 2.5 \times \frac{\rho_{NIR_1} - \rho_{red}}{\rho_{NIR_1} + 6 \times \rho_{red} - 7.5 \times \rho_{blue} + 1} \quad (1)$$

$$LSWI = \frac{\rho_{NIR_1} - \rho_{SWIR_1}}{\rho_{NIR_1} + \rho_{SWIR_1}} \quad (2)$$

where ρ_{NIR_1} , ρ_{red} , ρ_{blue} , and ρ_{LSWI} are the MOD09A1 surface reflectance for NIR_1 (841–876 nm), red (620–670 nm), blue (459–479 nm), and $SWIR_1$ (1628–1652 nm), respectively. A two-step gap-filling procedure was applied to gap-fill bad-quality observations within the time series of vegetation indices (Xiao et al., 2004a; Xiao et al., 2004b).

2.2. In-situ meteorological observations and CO₂ flux data

The AmeriFlux network consists of eddy covariance flux sites for monitoring the long-term ecosystem-scale exchange of carbon, energy, and water in North America (Baldocchi et al., 2001). Meteorological observations such as temperature, precipitation, and radiation are also collected at these sites.

We obtained all available 8-day Level 4 data of the AmeriFlux sites covering the conterminous U.S. from <http://ameriflux.lbl.gov/Pages/default.aspx> (Fig. 1). The Level 4 data included air temperature (T_{EC}), downward shortwave radiation (R_{EC}), and CO_2 flux data. This study used the standardized GPP (GPP_{EC}), which was partitioned from net ecosystem CO_2 exchange (NEE). By screening quality flags, only the most reliable observations were chosen for analysis. T_{EC} and R_{EC} from 37 non-crop sites (139 site-years) were used to evaluate and to adjust the NARR, if there were large biases. A total of 23 site-years of T_{EC} and R_{EC} and 40 site-years of GPP_{EC} from seven crop sites were used to validate the adjusted NARR and to evaluate the VPM-simulated GPP, respectively (Table 1). The crop sites were located in the Midwest U.S. corn and soybean belt, and were under different agricultural management practices. US-NE1 was a continuous irrigated maize site and US-NE2 was an irrigated maize/soybean rotation site. The other five sites were rainfed rotation sites. The detailed descriptions about these sites can be found in site specific publications (Griffis et al., 2005; Meyers and Hollinger, 2004; Verma et al., 2005).

[Insert Fig. 1 here]

[Insert Table 1 here]

It is important to mention that a direct comparison between the in-situ AmeriFlux observations and the NARR data without considering the differences of spatial scales might introduce some uncertainties. The in-situ observations can be affected by local environment conditions (terrain, hydrology, land cover etc.), while the NARR might be too coarse to delineate local environment variations. However, the AmeriFlux is currently the best available dataset providing high-quality and synchronized observation of radiation, temperature, water and carbon fluxes under standard protocols.

2.4. The Vegetation Photosynthesis Model (VPM)

The VPM is one PEM based on the conceptual partition of the light absorption by chlorophyll pigments and nonphotosynthetic vegetation (NPV such as branches, trunks, or senescent leaves) (Xiao et al., 2004a; Xiao et al., 2004b). The VPM defines the fPAR as the fraction of PAR absorbed by plant chlorophyll ($fPAR_{chl}$):

$$GPP = \varepsilon \times fPAR_{chl} \times PAR \quad (3)$$

$$fPAR_{chl} = EVI \quad (4)$$

$$\varepsilon = \varepsilon_0 \times T_{scalar} \times W_{scalar} \quad (5)$$

where PAR is calculated as $0.45 \times R$ (R , downward shortwave solar radiation); $fPAR_{chl}$ is equivalent to EVI; Light use efficiency, ε , is estimated as a function of the maximum light use efficiency (ε_0), temperature (T_{scalar}) and water condition (W_{scalar}). The ε_0 values of 3.12 g C MJ^{-1} for maize (Kalfas et al., 2011) and 1.75 g C MJ^{-1} for soybean (Wagle et al., 2015) were used in this study.

The effect of temperature scalar (T_{scalar}) on GPP is calculated using the equation from the Terrestrial Ecosystem Model (Raich et al., 1991):

$$T_{scalar} = \begin{cases} \frac{(T-T_{min})(T-T_{max})}{[(T-T_{min})(T-T_{max})]-(T-T_{opt})^2}, & T_{min} \leq T \leq T_{max} \\ 0, & T \leq T_{min}, T \geq T_{max} \end{cases} \quad (6)$$

where T is 8-day mean air temperature; T_{min} , T_{opt} , and T_{max} are minimum, optimum, and maximum temperatures for vegetation photosynthesis, respectively, and were set to 10°C , 28°C , 48°C for maize (Kalfas et al., 2011), and -1°C , 28°C , 50°C for soybean (Wagle et al., 2015).

The effect of water scalar (W_{scalar}) on GPP is calculated with LSWI:

$$W_{scalar} = \begin{cases} \frac{1+LSWI}{1+LSWI_{max}}, & LSWI > 0 \\ LSWI + LSWI_{max}, & LSWI \leq 0 \end{cases} \quad (7)$$

where $LSWI_{max}$ is the maximum LSWI during growing season.

This study used the VPM to simulate three sets of GPP_{VPM} : $GPP_{VPM(EC)}$, $GPP_{VPM(NARR)}$, and $GPP_{VPM(adjNARR)}$, using T and R from eddy flux sites (T_{EC} , R_{EC}), the NARR (T_{NARR} , R_{NARR}), and the adjusted NARR (T_{NARR} , $R_{adjNARR}$), respectively.

2.5. Statistical analysis

To quantify the differences between T_{NARR} and T_{EC} , R_{NARR} and R_{EC} , correlation coefficient (ρ), ratio of standard deviation (σ_{ratio}), bias, and root-mean-square-error (RMSE) were calculated for each non-crop site-year. The histogram of each statistics was summarized for all non-crop site-years to characterize the overall accuracy of T_{NARR} and R_{NARR} .

Mean Squared Error (MSE) was calculated for T_{NARR} and R_{NARR} of each site-year, and decomposed into three terms (Decker et al., 2012; Gupta et al., 2009), such that

$$MSE = 2\sigma_{NARR}\sigma_{EC} (1 - \rho) + (\sigma_{NARR} - \sigma_{EC})^2 + (\mu_{NARR} - \mu_{EC})^2 \quad (8)$$

where μ_{EC} and σ_{EC} are the mean and standard deviation for the in-situ observations, respectively. μ_{NARR} and σ_{NARR} are the mean and standard deviation for the NARR, respectively. The first, second, and third terms in equation (8) were represented in ternary diagrams to concisely visualized the contribution of correlation (ρ), consistency of variation (σ_{ratio}), and bias (bias and RMSE) to the overall disagreements between T_{NARR} and T_{EC} and between R_{NARR} and R_{EC} .

The simple linear regression between R_{EC} and R_{NARR} was also calculated for all non-crop site-years ($R_{EC} = \alpha \times R_{NARR}$). On the basis of the spatial pattern of regression coefficients (α), an empirical ratio-based adjustment was applied to R_{NARR} at the crop sites ($R_{adjNARR}$). Relative error (RE), RMSE, regression coefficient (α), and coefficient of determination (R^2) of the simple linear regression between R_{NARR} and R_{EC} , and $R_{adjNARR}$ and R_{EC} were obtained to quantify the adjustment performance.

This study implemented a top-down strategy to evaluate the impact of different climate inputs on GPP_{VPM} . First, the statistics factors described above were used to quantify how $GPP_{VPM(EC)}$, $GPP_{VPM(NARR)}$, and $GPP_{VPM(adjNARR)}$ matched GPP_{EC} for individual crops. Second, the similarities between GPP_{VPM} and GPP_{EC} across individual crop-sites were evaluated using Taylor diagrams. Taylor diagrams provide a statistical summary of the similarity of variability pattern (ρ), the agreement of the variability amplitudes (represented by the ratio of normalized standard deviation, $\bar{\sigma}_{ratio}$), and the centered RMSE between the modeled results and the observations (Gleckler et al., 2008; Taylor, 2001). In addition, annual mean RMSE of GPP_{VPM} was calculated for each crop site-year.

3. Results

3.1. Comparison of air temperature

T_{NARR} agreed well with T_{EC} for almost all non-crop site-years. T_{NARR} was significantly correlated with T_{EC} ($\rho > 0.95$ for 139 site-years, Fig. 2). In addition, T_{NARR} showed a similar amplitude of variation as in T_{EC} , as ~82 % of site-years had σ_{ratio} within ± 10 % error. T_{NARR} was mostly overestimated with a positive bias of 0.5–2.5 °C and a mean RMSE of 1.67 °C. The simple linear regression confirmed the good agreement between T_{NARR} and T_{EC} . T_{NARR} showed a strong linear regression with T_{EC} (α across 129 site-years was in a range of 1 ± 0.1 , $R^2 > 0.95$, $p < 0.001$). MSE was determined by both the bias and correlation, as the contribution of bias and correlation was over 0.8 at 86 % of the site-years (Fig. 3).

[Insert Fig. 2 here]

[Insert Fig. 3 here]

T_{NARR} was also relatively accurate at the crop sites. The simple linear regression indicated that T_{NARR} agreed well with T_{EC} for all crop site-years ($\alpha = 1.04$, $RE = 11.6\%$, $RMSE = 1.4\text{ }^{\circ}C$, $R^2 = 0.99$, Fig. 4). T_{NARR} accounted for over 98 % of the seasonal dynamics of T_{EC} for individual crop sites on annual scale (Table 2). α varied from 1.0 to 1.1 among the crop sites. RE and $RMSE$ were $-1.4\text{--}7.3\%$ and $1.2\text{--}1.7\text{ }^{\circ}C$, respectively. Considering the relatively high accuracy at non-crop and crop site-years, the 8-day T_{NARR} was used as the VPM input without any correction.

[Insert Fig. 4 here]

[Insert Table 2 here]

3.2. Comparison of downward shortwave radiation

R_{NARR} was well correlated with R_{EC} ($\rho > 0.9$ at 94 % of the non-crop site-years, Fig. 5). However, it was overestimated with $\sigma_{ratio} > 1.1$ at 67 % of the site-years. The bias was positive across all site-years on an average of $3.55\text{ MJ m}^{-2}\text{ day}^{-1}$. 60 % of the site-years had a $RMSE$ of $3\text{--}5\text{ MJ m}^{-2}\text{ day}^{-1}$. The bias was the dominant contributor to MSE (Fig. 6). The contribution of bias was > 0.5 at 133 of 139 site-years, indicating the disagreement between R_{NARR} and R_{EC} was systematic.

[Insert Fig. 5 here]

[Insert Fig. 6 here]

R_{NARR} showed a significant linear regression with R_{EC} at each non-crop site-year (Fig. 5). However, α was quite variable ($0.63\text{--}0.95$). α slightly decreased with the latitude increasing or the longitude decreasing (Fig. 7). α was more stable within the longitude range of $85\text{--}100\text{ }^{\circ}W$ than it was across $40\text{--}47.5\text{ }^{\circ}N$ for the region covering the crop sites (Fig. 7 highlighted in gray).

Thus, the median of α values (0.81) within the longitude of 85–100 °W was used as a ratio to adjust the bias of R_{NARR} at the crop sites.

[Insert Fig. 7 here]

The adjustment substantially reduced the bias of R_{NARR} at the crop sites (Fig. 8). R_{NARR} was overestimated by 28.2 % on average. $R_{adjNARR}$ evenly distributed along 1:1 line and RMSE was reduced to $1.7 \text{ MJ}^{-2} \text{ day}^{-1}$.

[Insert Fig. 8 here]

R_{NARR} explained ~ 90 % of the variations of R_{EC} across each crop site (Table 2). Similar to the non-crop sites, R_{NARR} was strongly overestimated ($RE > 22 \%$) at the crop sites. The annual RMSE varied from $3.8 \text{ MJ m}^{-2} \text{ day}^{-1}$ to $4.9 \text{ MJ m}^{-2} \text{ day}^{-1}$. After the adjustment, α was close to 1, and RE and RMSE of $R_{adjNARR}$ decreased to -2.5–3 % and $1.6\text{--}2 \text{ MJ m}^{-2} \text{ day}^{-1}$, respectively.

3.3 Comparison of VPM-based (GPP_{VPM}) and the flux tower-based (GPP_{EC}) estimates

The seasonal dynamics of $GPP_{VPM(EC)}$, $GPP_{VPM(NARR)}$, and $GPP_{VPM(adjNARR)}$ corresponded well with GPP_{EC} for both maize and soybean (Fig. 9). At the leaf-on stage during late-May to June, GPP_{EC} started to exceed $1 \text{ g C m}^{-2} \text{ day}^{-1}$ and GPP_{VPM} also rose rapidly, and both reached a maximum at the peak growing season during late-July to early-August. After the crops matured and approached the harvest date in September, both GPP_{EC} and GPP_{VPM} began to decrease and were lower than $1 \text{ g C m}^{-2} \text{ day}^{-1}$.

[Insert Fig. 9 here]

The relationships between GPP_{VPM} and GPP_{EC} for individual crop types were evaluated through simple linear regression models (Fig. 10). For irrigated and rainfed maize, both $GPP_{VPM(EC)}$ and $GPP_{VPM(adjNARR)}$ agreed well with GPP_{EC} ; but $GPP_{VPM(NARR)}$ was overestimated

due to the positive bias of R_{NARR} (Figs. 10a and 10b). $GPP_{VPM(EC)}$ accounted for 89 % of the variations of GPP_{EC} . $GPP_{VPM(NARR)}$ was also correlated well with GPP_{EC} , but it was overestimated by 18.4 % and 29.6 % for irrigated and rainfed maize, respectively. After adjusting R_{NARR} , α , RE, and RMSE for $GPP_{VPM(adjNARR)}$ were close to those of $GPP_{VPM(EC)}$. For irrigated and rainfed soybean, $GPP_{VPM(EC)}$ and $GPP_{VPM(adjNARR)}$ estimated GPP reasonably well with an underestimation < -10 % (Figs. 10c and 10d). $GPP_{VPM(NARR)}$ over-predicted GPP_{EC} by ~13 % for irrigated and rainfed soybean.

[Insert Fig. 10 here]

The relationships between GPP_{VPM} and GPP_{EC} were further evaluated for maize through individual crop-sites and individual site-years (Fig. 11a, 11b, and Table 3). $GPP_{VPM(EC)}$ and $GPP_{VPM(adjNARR)}$ showed reliable GPP estimates for the irrigated and rainfed maize across the sites (Figs. 11a and 11b). Most sites had similar patterns and amplitudes of variability between $GPP_{VPM(EC)}$ and GPP_{EC} ($1 < \bar{\sigma}_{ratio} < 1.05$ and $0.95 < \rho < 0.98$, Fig. 11a) with low annual mean RMSEs (ca. 1.5–2.4 g C m⁻² day⁻¹, Table 3). $GPP_{VPM(EC)}$ at RO1 and Bo1 didn't appear to adequately capture the amplitudes of variability of GPP_{EC} ($\bar{\sigma}_{ratio}$ = 0.7 and 1.3) as indicated by relatively low ρ (0.92 and 0.82) and high RMSE (3.2 g C m⁻² day⁻¹ and 4.9 g C m⁻² day⁻¹). The discrepancies were due to the underestimation of $GPP_{VPM(EC)}$ during the peak growing season at RO1 and the significant overestimation of $GPP_{VPM(EC)}$ after the peak growing season at Bo1 (Fig. 9). $GPP_{VPM(NARR)}$ simulated the phasing and timing of GPP_{EC} well (ρ was ca. 0.93–0.98). The RMSE of $GPP_{VPM(NARR)}$ (ca. 4.2–6.4 g C m⁻² day⁻¹) was significantly higher than that of $GPP_{VPM(EC)}$ at most sites, indicating an overestimation caused by the NARR. The adjustment of R_{NARR} resulted in similar patterns of $GPP_{VPM(adjNARR)}$ and $GPP_{VPM(EC)}$ at all sites, with a slight increase of RMSE (ca. 1.6–3.1 g C m⁻² day⁻¹, Fig 11b and Table 3).

[Insert Fig. 11 here]

[Insert Table 5 here]

The relationships between GPP_{VPM} and GPP_{EC} were also evaluated for soybean through individual crop-sites and individual site-years (Fig. 11c, 11d, and Table 3). $GPP_{VPM(EC)}$ and $GPP_{VPM(adjNARR)}$ matched GPP_{EC} reasonably well. The variability of $GPP_{VPM(EC)}$ was similar to that of GPP_{EC} ($0.83 < \rho < 0.93$, Fig. 11c). NE2 and RO1 had a good agreement between $GPP_{VPM(EC)}$ and GPP_{EC} , as $\bar{\sigma}_{ratio}$ was close to 1 showing a low RMSE ($1.4\text{--}2.0 \text{ g C m}^{-2} \text{ day}^{-1}$, Table 3). At other sites (NE3, IB1, and Bo1), $GPP_{VPM(EC)}$ underestimated the variability of GPP_{EC} ($0.85 < \bar{\sigma}_{ratio} < 0.9$, Fig. 11c) with a high RMSE ($2.0\text{--}2.6 \text{ g C m}^{-2} \text{ day}^{-1}$). $GPP_{VPM(NARR)}$ correlated well with GPP_{EC} ($0.9 < \rho < 0.94$). However, the $\bar{\sigma}_{ratio}$ of $GPP_{VPM(NARR)}$ was larger than that of $GPP_{VPM(EC)}$ caused by the positive bias of R_{NARR} . After adjusting the bias of R_{NARR} , $GPP_{VPM(adjNARR)}$ matched GPP_{EC} better than did $GPP_{VPM(NARR)}$ (Fig. 11d).

4. Discussion

4.1. Uncertainties of the NARR air temperature

T_{NARR} has been assumed to be relatively accurate in the studies of drought monitoring and the response of vegetation to climate change (Karnauskas et al., 2008; Karnieli et al., 2010; Wang et al., 2011). In this study, the 8-day T_{NARR} was mostly overestimated with a mean bias of $0.62 \text{ }^{\circ}\text{C}$. This was consistent with the previous finding that T_{NARR} was biased warm at monthly intervals (Jiang and Yang, 2012). In general, the 8-day T_{NARR} agreed well with the in-situ observations across non-crop and crop site-years with the mean RMSE of $1.67 \text{ }^{\circ}\text{C}$ and $1.4 \text{ }^{\circ}\text{C}$, respectively, showing relatively higher accuracy than other global reanalysis datasets (DAO, ECMWF, NCEP, MERRA) investigated by Zhao et al. (2006) and Decker et al. (2012).

4.2. Uncertainties of the NARR downward shortwave radiation

This study made an assumption that R_{EC} were ground truth. However, the errors or uncertainties associated with in-situ radiation observations also contributed to the differences between R_{NARR} and R_{EC} . R_{EC} at the AmeriFlux is measured by different pyranometers. The errors from pyranometers including instrument deployment and maintenance (leveling and shading) and sensor response errors such as thermal offset (Bush et al., 2000; Reda et al., 2005) determined the errors of R_{EC} . The errors of R_{EC} are subtle compared with the R_{NARR} biases, but one shouldn't neglected their impacts considering the significant decay of long-term sensor stability (Stanhill and Cohen, 2001).

A number of studies have evaluated the monthly R_{NARR} at individual sites. Walsh (2009) evaluated the montly R_{NARR} at the Alaska Barrow site and found that it had a lower bias ($2.6 \text{ MJ m}^{-2} \text{ day}^{-1}$) than did NCEP/NCAR ($3.7 \text{ MJ m}^{-2} \text{ day}^{-1}$). Kennedy et al. (2011) concluded that the bias of monthly R_{NARR} varied with sky conditions at the Atmospheric Radiation Measurement Program (ARM) Southern Great Plain (SGP) site. Markovic et al (2009) reported that a systematic bias of monthly R_{NARR} in summer ($5.3 \text{ MJ m}^{-2} \text{ day}^{-1}$) was larger than that in winter ($2.5 \text{ MJ m}^{-2} \text{ day}^{-1}$). These evaluations implied that R_{NARR} had a large span of positive biases. However, their results cannot represent the overall accuracy of R_{NARR} at continental scale using limited sites. A recent study did a large-scale assessment of monthly R_{NARR} using 24 FLUXNET sites showing that R_{NARR} exhibited a positive bias of $3.2 \text{ MJ m}^{-2} \text{ day}^{-1}$ (Zhao et al., 2013a). The ideal temporal interval of climatic drivers for ecological models should be finer, i.e. hourly, daily, or weekly intervals, to demonstrate the diurnal or seasonal dynamics of carbon and energy fluxes (Abatzoglou, 2013; Huntzinger et al., 2013; Wei et al., 2013). Thus, we evaluated the accuracy of R_{NARR} at 8-day intervals and regional scale using all available AmeriFlux sites. The 8-day

R_{NARR} well represented the seasonal dynamics of 8-day R_{EC} . Similar to monthly R_{NARR} , the bias of the 8-day R_{NARR} was positive and systematic with a large range across the U.S.. The systematic overestimation of R_{NARR} is mainly caused by the insufficient simulation of light extinction caused by clouds, aerosols, and water vapor in the radiative transferring models (Kennedy et al., 2011; Markovic et al., 2009; Zhao et al., 2013a), and other topographical factors (i.e. elevation, slope, and aspect) (Schroeder et al., 2009; Zhao et al., 2013a).

Empirical or semi-empirical approaches are applied to correct the bias of R_{NARR} . The empirical approach develops the linear statistical regression model between the reanalysis and in-situ observations, then applies the model to other locations (Feng et al., 2007; Qian et al., 2006; Xiao et al., 2014). The empirical approach ignores the spatio-temporal variations in the R_{NARR} bias. Some studies developed the semi-empirical approach to account the impacts of clouds and topographical factors in the regression models (Schroeder et al., 2009; Zhao et al., 2013a). We followed the empirical approach to calibrate R_{NARR} , and meanwhile considered the spatial variation of regression models.

Simply estimating PAR as a constant ratio of R_{NARR} can introduce uncertainties to PAR. Theoretically, the band range of downward shortwave radiation (0.3–2.8 μm) doesn't match that of PAR (0.4–0.7 μm) (Sakamoto et al., 2011). Moreover, the ratio of PAR to downward shortwave radiation is not constant, as it temporally changes with the local weather condition (Gonzalez and Calbo, 2002; Jacovides et al., 2004; Papaioannou et al., 1993). Surface PAR datasets, such as the satellite-derived Global Land Surface Satellite (GLASS), might be an alternative PAR input for the regional and global ecological modeling (Cai et al., 2014; Eck and Dye, 1991; Frouin and Pinker, 1995; Jin et al., 2013; Pinker et al., 2010; Rubio et al., 2005; Zhao et al., 2013b).

4.3 Sensitivity of PEMs to various climate inputs

All analyses about the sensitivity of PEMs to climate inputs were focused on the PEM of the standard MODIS GPP product —the MODIS-PSN (Heinsch et al., 2006; Zhang et al., 2007; Zhao et al., 2006). These studies found that radiation, air temperature, and vapor pressure deficit (VPD) of the global reanalysis data were largely biased, and introduced significant errors to the standard MODIS GPP product. For instance, Zhao et al. (2006) found that the MODIS GPP showed significant differences when driven by DAO, NCAR, and ECMWF ($> 20 \text{ Pg C yr}^{-1}$). Heinsch et al. (2006) collected 38 site-years of GPP_{EC} from 15 AmeriFlux sites to evaluate the accuracy of MODIS GPP driven by DAO and in-situ meteorology, and annual GPP derived from DAO was 23 % higher than GPP_{EC} and the RE of the GPP derived from DAO was much larger than that of the GPP derived from the in-situ meteorology. Note that these evaluations were conducted at monthly or longer intervals. Analyses on finer temporal scales such as weekly interval are needed in order to accurately evaluate the seasonal dynamics of the uncertainties of PEMs to climate data. Thus, we focused on quantifying the uncertainties of GPP_{VPM} to in-situ and NARR climate data at 8-day interval. The 8-day GPP_{VPM} driven by the in-situ meteorology, original and adjusted NARR data traced over 83–98 % of GPP_{EC} variations for individual site-years, confirming their capabilities to simulate the response of crop photosynthesis to the environment change (i.e. light, temperature, and water), and tracked the phenological phases well (i.e. leaf-on and leaf-off stages). Similar to the MODIS-PSN, climate inputs had a strong impact on the VPM for cropland GPP estimates. $\text{GPP}_{\text{VPM}(\text{EC})}$ well estimated GPP_{EC} for individual crops, sites, and site-years. $\text{GPP}_{\text{VPM}(\text{NARR})}$ significantly overestimated GPP_{EC} as R_{NARR} was positively biased. This study addressed two climate inputs of air temperature and downward shortwave radiation for the VPM. The accuracies of other climate variables (VPD, precipitation etc.) in

reanalysis products might be more variable (Decker et al., 2012). Therefore, more uncertainties might be introduced to the PEMs that are driven by multiple climate variables.

4.4 Challenges in comparing GPP_{VPM} with GPP_{EC}

In one study like ours using GPP_{EC} to validate or constrain the GPP estimates from PEMs, two assumptions are often made: (1) GPP_{EC} is assumed to be accurate as the ground truth and (2) The eddy flux tower footprint is approximately equivalent to the image pixel. The uncertainties associated with these two assumptions, however, can contribute to the discrepancies between the PEM-based GPP estimates (GPP_{VPM} in this study) and GPP_{EC} .

There are a number of errors or uncertainties (random and systematic) from eddy covariance measurements. Random errors are attributed to the stochastic nature of turbulence, sampling errors, instrument system, and variations in the flux footprint (Richardson et al., 2012). Systematic errors arise from the combination of the unmet underlying theoretical assumptions, instrument calibration, and data processing techniques (Falge et al., 2001; Papale et al., 2006; Richardson et al., 2012). Furthermore, the eddy covariance provides direct measurement of NEE and GPP_{EC} is estimated as the difference between NEE and ecosystem respiration (R_{eco}) using flux-partitioning approaches, which may also introduce large uncertainties in GPP_{EC} (Desai et al., 2008; Reichstein et al., 2005; Stoy et al., 2006). For example, Desai et al. (2008) found annual GPP_{EC} varied $\sim 100 \text{ g C m}^{-2} \text{ year}^{-1}$ among 23 partitioning methods. Thus, more efforts are needed to improve partitioning NEE into its gross components to help validate GPP in PEMs and other land surface models (Baldocchi et al., 2015).

The second assumption is questionable in heterogeneous landscapes. Limited by data availability, most PEMs are performed on 1 km spatial resolution of satellite images and might

not represent the crop fields that towers are located in due to the mixed signals from other sub-pixel components. In this study, an in-situ landscape analysis showed that the heterogeneity of 500 m MODIS pixels was much improved over that of 1 km MODIS pixels at seven crop sites (Fig. S1). 500 m MODIS pixels were mainly covered by the crop fields that the towers measured except US-RO3 and US-Bo1. Even though the uncertainties of the GPP comparison caused by heterogeneous landscapes were diminished to some extent using the 500 m MODIS data in this study, further evaluations using high resolution images along with the downscaling techniques are required for implementing PEMs, especially at heterogeneous landscapes.

5. Conclusion

This study evaluated the uncertainties of the NARR surface meteorology and quantified the sensitivity of the VPM to the in-situ and NARR climate inputs at seven AmeriFlux crop eddy flux sites. Our results indicated that the bias of NARR resulted in considerable uncertainties in cropland GPP estimates. The 8-day NARR air temperature matched well with in-situ observations, but the NARR downward shortwave radiation showed large positive bias and led to the overestimation of GPP_{VPM} . An empirical correction of the NARR radiation improved the model performance.

The findings of this study confirm the good performance of the VPM on estimating maize and soybean GPP as long as meteorological inputs are accurate, and imply that the capability of the satellite-based PEMs for regional productivity monitoring at heterogeneous landscapes would be enhanced if the radiation of the regional reanalysis product can be improved to resolve the impacts of cloud cover and terrain. The proposed method to correct NARR radiation is limited to the crop sites in this study, and might not be applicable for other regions

due to the large spatial variations of the NARR radiation bias. In addition to the meteorological data, further research is required to address the uncertainties of the PEM-based GPP estimates caused by other model inputs such as satellite data.

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Appendix A. Supplementary data

Supplementary data associated with this article can be found in the online version, at

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List of Tables

Table 1. A summary description of the AmeriFlux eddy flux crop sites.

Site ID	Latitude	Longitude	Years of T_{EC}^a , R_{EC}^b	Crop type	Years of GPP_{EC}^c
US-NE1	-96.4766	41.165	2002-2012	Irrigated maize	2002–2005, 2007–2012
US-NE2	-96.4701	41.1649	2001-2012	Irrigated maize	2001, 2003, 2005, 2007, 2009–2012
				Irrigated soybean	2002, 2004, 2006, 2008
US-NE3	-96.4396	41.1797	2001-2012	Rainfed maize	2001, 2003, 2005, 2009, 2011
				Rainfed soybean	2002, 2004, 2008, 2010
US-RO1	-93.0898	44.7143	2004-2006	Rainfed maize	2005
				Rainfed soybean	2006
US-RO3	-93.0893	44.7217	2004-2006	Rainfed maize	2005
US-IB1	-88.2904	40.0062	2006-2007	Rainfed maize	2006
				Rainfed soybean	2007
US-Bo1	-88.2227	41.8593	2001-2006	Rainfed maize	2005
				Rainfed soybean	2002, 2004, 2006

^a and ^b air temperature and downward shortwave radiation observed from the AmeriFlux eddy flux sites

^c 8-day Level-4 GPP estimates from the AmeriFlux eddy flux sites

Table 2. Statistics of the comparison of the 8-day NARR air temperature, original, and adjusted downward shortwave radiation with the AmeriFlux observations for the individual crop sites.

Site ID	T_{EC}^a vs. T_{NARR}^b				R_{EC}^c vs. R_{NARR}^d				R_{EC} vs. $R_{adjNARR}^e$		
	RE	RMSE	α	R^2	RE	RMSE	α	R^2	RE	RMSE	α
US-NE1/2/3	7.3±5.9	1.7±0.55	1.1±0.04	0.98	21.8±2.6	3.9±0.4	1.2±0.03	0.9±0.03	-2.5±2.1	1.6±0.02	0.97±0.02
US-RO1/3	-1.4±5.2	1.6±0.33	1.0±0.01	0.99±0.01	26.5±2.2	4.2±0.2	1.2±0.02	0.91±0.03	1.2±1.8	1.6±0.22	0.99±0.02
US-IB1	4.6±0.1	1.3±0.12	1.1±0.02	0.99	28.6±1.5	4.9±0.26	1.3	0.89±0.05	2.9±1.2	1.8±0.23	1.01
US-Bo1	7.1±2.0	1.2±0.17	1.0±0.01	0.99	22.7±7.6	3.8±0.78	1.2±0.08	0.9±0.05	-1.9±6.1	2.0±0.40	0.96±0.06

^a and ^b air temperature of the AmeriFlux and NARR data (°C)

^c, ^d and ^e downward shortwave radiation of the AmeriFlux, the NARR before and after adjustment ($MJ\ m^{-2}\ day^{-1}$)

Table 3. A summary of the performances of the VPM driven by three sets of climate inputs at the crop sites.

Site ID	Crop type	GPP _{VPM(EC)} ^a			GPP _{VPM(NARR)} ^b			GPP _{VPM(adjNARR)} ^c		
		ρ	$\bar{\sigma}_{ratio}$	RMSE	ρ	$\bar{\sigma}_{ratio}$	RMSE	ρ	$\bar{\sigma}_{ratio}$	RMSE
US-NE1	Irrigated maize	0.95	1.03	2.4±0.7	0.93	1.25	5.0±1.1	0.93	1.00	2.7±0.6
US-NE2	Irrigated maize	0.96	1.02	2.2±0.8	0.95	1.24	4.8±0.6	0.95	0.99	2.3±0.6
	Irrigate soybean	0.91	1.00	2.0±0.6	0.91	1.18	2.3±0.7	0.90	0.94	2.0±0.6
US-NE3	Rainfed maize	0.95	1.00	2.1±0.5	0.95	1.26	4.2±1.3	0.95	1.01	2.2±0.5
	Rainfed soybean	0.91	0.89	2.6±0.4	0.91	1.05	2.2±0.3	0.90	0.84	2.8±0.5
US-RO1	Rainfed maize	0.92	0.7	3.2	0.93	0.86	4.8	0.93	0.69	3.1
	Rainfed soybean	0.92	1.06	1.4	0.92	1.27	3.4	0.92	1.02	1.5
US-RO3	Rainfed maize	0.98	1.03	1.5	0.97	1.23	4.3	0.98	0.98	1.6
US-IB1	Rainfed maize	0.97	1.06	1.5	0.96	1.45	6.4	0.96	1.16	3.1
	Rainfed soybean	0.83	0.88	2.3	0.93	1.11	4.1	0.93	0.89	1.9
US-Bo1	Rainfed maize	0.82	1.31	4.9	0.78	1.66	8.6	0.78	1.33	5.5
	Rainfed soybean	0.93	0.86	2.0	0.91	1.03	2.5	0.91	0.83	2.3

^{a, b, c} the VPM-based GPP estimates from in-situ, original and adjusted NARR climate data

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Fig. 1. Location of the AmeriFlux eddy flux sites. Circles denote the non-crop sites for accuracy assessment of the NARR and stars denote the crop sites used to evaluate the VPM-based GPP estimates. The base map is the 2013 Cropland Data Layer (CDL) from the National Agricultural Statistics Service (NASS).

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Fig. 6. Contributions of correlation (ρ), consistency of variation (σ_{ratio}), and bias to the Mean Squared Error (MSE) for the 8-day NARR downward shortwave radiation (R_{NARR}) across the non-crop site-years.

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737 the irrigated and rainfed soybean. The locations of the heads and tails of arrows quantify how
738 GPP_{VPM} matches with GPP_{EC} , and the arrows show how the **agreement** of GPP_{VPM} with **GPP_{EC}**
739 changes using different climate inputs. The distance to the origin is the ratio of the standard
740 deviations of GPP_{VPM} and GPP_{EC} (Normalized standard deviation, $\bar{\sigma}_{ratio}$). The azimuthal angle is

741 the correlation (ρ) showing the similarity of variation patterns between GPP_{VPM} and GPP_{EC} . The
742 most ideal GPP_{VPM} estimate is the point “observed” with $\bar{\sigma}_{ratio} = 1$ and $\rho = 1$.

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